

Photo-oxidation of aromatic hydrocarbons produces low-volatility organic compounds

Mingyi Wang^{1,2}, Dexian Chen^{1,3}, Mao Xiao⁴, Qing Ye^{1,2}, Dominik Stolzenburg⁵, Victoria Hofbauer^{1,2}, Penglin Ye⁶, Alexander L. Vogel⁷, Roy L. Mauldin III^{8,1,2}, Antonio Amorim⁹, Andrea Baccarini⁴, Bernhard Baumgartner⁵, Sophia Brilke⁵, Lubna Dada¹⁰, António Dias⁹, Jonathan Duplissy^{10,11}, Henning Finkenzeller¹², Olga Garmash¹⁰, Xu-Cheng He¹⁰, Christopher R. Hoyle^{4,13}, Changhyuk Kim^{14,15}, Alexander Kvashnin¹⁶, Katrianne Lehtipalo^{10,17}, Lukas Fischer¹⁸, Ugo Molteni⁴, Tuukka Petäjä¹⁰, Veronika Pospisilova⁴, Lauriane L. J. Quéléver¹⁰, Matti Rissanen¹⁹, Mario Simon⁷, Christian Tauber⁵, António Tomé²⁰, Andrea C. Wagner^{7,12}, Lena Weitz⁷, Rainer Volkamer¹², Paul M. Winkler⁵, Jasper Kirkby^{21,7}, Douglas R. Worsnop^{10,6}, Markku Kulmala^{10,11,22,23}, Urs Baltensperger⁴, Josef Dommen⁴, Imad El-Haddad⁴, and Neil M. Donahue^{1,2,3,24,*}

¹Center for Atmospheric Particle Studies, Carnegie Mellon University, Pittsburgh, PA 15213, USA;
²Department of Chemistry, Carnegie Mellon University, Pittsburgh, PA, 15213, USA; ³Department of Chemical Engineering, Carnegie Mellon University, Pittsburgh, PA, 15213, USA; ⁴Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, 5232 Villigen, Switzerland; ⁵Faculty of Physics, University of Vienna, 1090 Vienna, Austria; ⁶Aerodyne Research, Inc., Billerica, Massachusetts 01821, USA; ⁷Institute for Atmospheric and Environmental Sciences, Goethe University Frankfurt, 60438, Frankfurt am Main, Germany; ⁸Department of Oceanic and Atmospheric Science, University of Colorado Boulder, Boulder, CO 80309, USA; ⁹CENTRA and FCUL, University of Lisbon, 1749-016 Lisbon, Portugal; ¹⁰Institute for Atmospheric and Earth System Research (INAR), University of Helsinki, 00014 Helsinki, Finland; ¹¹Helsinki Institute of Physics, University of Helsinki, 00014 Helsinki, Finland; ¹²Department of Chemistry & CIRES, University of Colorado Boulder, Boulder, CO 80309, USA; ¹³Institute for Atmospheric and Climate Science, ETH Zurich, 8092 Zurich, Switzerland; ¹⁴Division of Chemistry and Chemical Engineering, California Institute of Technology, Pasadena, CA 91125, USA; ¹⁵Department of Environmental Engineering, Pusan National University, Busan, 46241, Republic of Korea; ¹⁶Lebedev Physical Institute, 119991, Moscow, Russia; ¹⁷Finnish meteorological Institute, Erik Palménin aukio 1, 00560 Helsinki, Finland; ¹⁸Institute for Ion Physics and Applied Physics, University of Innsbruck, 6020 Innsbruck, Austria; ¹⁹Aerosol Physics Laboratory, Physics Unit, Tampere University, P.O.Box 1001, Tampere, Finland; ²⁰IDL-University of Beira Interior, Covilhã, Portugal; ²¹CERN, 1211 Geneva, Switzerland; ²²Joint International Research Laboratory of Atmospheric and Earth System Sciences, Nanjing University, Nanjing, China; ²³Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering, Beijing University of Chemical Technology, Beijing, China; and ²⁴Department of Engineering and Public Policy, Carnegie Mellon University, Pittsburgh, PA, 15213, USA.

*To whom correspondence should be addressed (Email: nmd@andrew.cmu.edu).

37 ABSTRACT

38 To better understand the role of aromatic hydrocarbons in new-particle formation, we measured
39 the particle-phase abundance and volatility of oxidation products following the reaction of
40 aromatic hydrocarbons with OH radicals. For this we used thermal desorption in an iodide-
41 adduct Time-of-Flight Chemical-Ionization Mass Spectrometer equipped with a Filter Inlet for
42 Gases and AEROsols (FIGAERO-ToF-CIMS). The particle-phase volatility measurements
43 confirm that oxidation products of toluene and naphthalene can contribute to the initial growth
44 of newly formed particles. Toluene-derived (C_7) oxidation products have a similar volatility
45 distribution to that of α -pinene-derived (C_{10}) oxidation products, while naphthalene-derived
46 (C_{10}) oxidation products are much less volatile than those from toluene or α -pinene; they are
47 thus stronger contributors to growth. Rapid progression through multiple generations of
48 oxidation is more pronounced in toluene and naphthalene than in α -pinene, resulting in more
49 oxidation but also favoring functional groups with much lower volatility per added oxygen
50 atom, such as hydroxyl and carboxylic groups instead of hydroperoxide groups. Under
51 conditions typical of polluted urban settings, naphthalene may well contribute to nucleation
52 and the growth of the smallest particles, whereas the more abundant alkyl benzenes may
53 overtake naphthalene once the particles have grown beyond the point where the Kelvin effect
54 strongly influences the condensation driving force.

55

56 INTRODUCTION

57 Organic condensation from the gas phase to the particle phase contributes to particle growth
58 and causes secondary organic aerosol (SOA) formation¹. SOA in turn may comprise well over
59 half of the total fine-particle mass in the lower atmosphere, contributing to and possibly
60 dominating health effects include more than 4 million premature deaths per year^{2,3}. However,
61 to what extent organic condensation affects the initial growth of nucleated particles remains in
62 doubt, and the mechanism forming "condensable" organic vapors is even less certain⁴. The
63 growth rate of freshly nucleated particles is crucial to their survival⁵, which in turn governs
64 the connection between particle nucleation and aerosol-cloud interactions, one of the most
65 uncertain elements of the climate system⁶.

66 Recent results from the "Cosmics Leaving OUtdoor Droplets" (CLOUD) experiment show

that organic vapors emitted by trees can produce abundant aerosol particles through nucleation of ultra-low volatility organic compounds (ULVOCs) even when there is no sulfuric acid ⁷. Tröstl *et al.* ¹, Lehtipalo *et al.* ⁸ and Stolzenburg *et al.* ⁹ have constrained the role of pure biogenic organic vapors in the initial growth of nucleated particles via laboratory experiments free from inorganic acids and bases or nitrogen oxides (NO_x). Combined experimental data and model simulations show that biogenic organic vapors driving initial growth up to ~2 nm are extremely low-volatility organic compounds (ELVOCs). As particles grow and the Kelvin effect diminishes, subsequent growth is taken over by more abundant low-volatility organic compounds (LVOCs) with slightly higher volatility ¹⁰.

Ozonolysis of terpenes, followed by peroxy-radical autoxidation via a succession of internal H-atom transfer reactions, produces ULVOCs and ELVOCs within a few seconds after the initial attack by ozone ¹¹. This drives new-particle formation in regions dominated by biogenic emissions ¹². However, these processes can be suppressed substantially in polluted airmasses, where nitrogen oxides (NO_x) can perturb peroxy-radical chemistry and suppress ULVOC and ELVOC formation ¹³. Further, even under optimal conditions, those compounds comprise only < 5% of the terpene oxidation products ¹⁴. Though first-generation autoxidation plays a major role, the (large) balance of “traditional” product vapors cannot be neglected; those vapors will undergo further oxidation, which has been shown to increase yields of SOA ¹⁵ and also to form more ELVOCs via subsequent autoxidation ¹⁶.

Pollution contains much more than NO_x . In the urban atmosphere, aromatic hydrocarbons can account for up to 60 % of the total VOCs ^{17,18}. Among them, small monocyclic aromatics are mostly emitted from anthropogenic sources, such as fossil-fuel combustion and solvent evaporation, with toluene being one of the most abundant species ^{17,18}. Polycyclic aromatic hydrocarbons mainly come from incomplete combustion such as heavy-duty diesel vehicles, biomass burning and meat cooking ^{19–21}; formation of naphthalene is often favored among PAHs during these combustion processes ²⁰. Aromatic hydrocarbons are believed to play a critical role in the formation of both tropospheric ozone and secondary organic aerosols (SOA) ^{22–25}. Laboratory studies have also reported highly oxygenated product formation from OH-triggered oxidation of aromatic hydrocarbons, highlighting their potential contribution to nucleation and initial particle growth in the anthropogenic environment ^{26,27}.

97 Despite the significance of aromatic systems to the urban atmosphere, the volatility profile
98 of oxidation products from aromatic precursors has yet to be established. The contribution of
99 aromatic-derived oxidation products to urban new-particle formation has so far remained
100 ambiguous. Here, we extend previous CLOUD experiments on biogenic organic vapors to add
101 toluene and naphthalene as examples of anthropogenic organic vapors, comparing the
102 formation pathways of oxidized organics as well as their behaviors on contributing to the
103 growth of newly formed particles in the urban environment.

104

105 MATERIALS AND METHODS

106 We conducted experiments involving aromatic precursors under typical urban conditions in the
107 “Cosmics Leaving OUtdoor Droplets” (CLOUD) chamber. The CLOUD facility was
108 constructed to measure nucleation and growth rates of inorganic and organic vapors under
109 highly controlled, ultra-clean conditions. The 26.1 m³ stainless steel chamber is kept
110 scrupulously clean, with background contaminants (H₂SO₄, organics, amines) < 10⁵ cm⁻³. It
111 can be exposed to ultraviolet (UV) light via a fiber-optic system to generate OH radicals
112 uniformly via ozone photolysis, initiating subsequent photochemical reactions. In addition, the
113 Xe arc lights with UVA at 385 nm from a set of 400 W LED lights are augmented to photolyze
114 NO₂ to NO. The chamber temperature can be controlled from 183 to 373 K ²⁸, and relative
115 humidity (RH) from < 0.5% to 101%, enabling experiments over a wide range of tropospheric
116 conditions.

117 **Experiment conditions.** Results presented here are from the CLOUD-11 campaign in Fall
118 2016. To simulate urban environments, we fixed the temperature at 20 °C, relative humidity at
119 60%, ozone concentration at ~ 40 ppbv, NH₃ concentration at ~ 500 pptv, SO₂ concentration
120 at ~ 2 ppbv, and NO₂ concentration at ~ 1–2 ppbv. We conducted experiments with different
121 aromatic precursor concentrations spanning urban conditions – 8 ppbv and 40 ppbv for toluene
122 and 2 ppbv and 15 ppbv for naphthalene. We formed OH radicals via ozone photolysis and the
123 subsequent reaction of O(¹D) with water vapor.

124 **Chemical ionization mass spectrometer.** We measured the gas-phase and particle-phase
125 composition as well as particle-phase volatility of oxidation products via thermal desorption
126 using an iodide-adduct Time-of-Flight Chemical Ionization Mass Spectrometer equipped with

127 a Filter Inlet for Gases and AEROSols (FIGAERO-ToF-CIMS). Iodide-adduct chemical
128 ionization is well suited for measuring a suite of oxygenated organics with minimal
129 fragmentation ²⁹. The FIGAERO is a manifold inlet for the CIMS with two operating modes
130 ³⁰. In the first mode, gases are directly sampled into a turbulent flow ion-molecule reactor while
131 particles are concurrently collected on a PTFE filter via a separate dedicated port. In the second
132 mode, the filter is automatically moved into a pure N₂ gas stream flowing into the ion-molecule
133 reactor while the N₂ is heated to evaporate the particles via temperature programmed desorption.
134 Analytes are then chemically ionized and extracted into a long ToF-MS, achieving sensitivities
135 below 10⁶ cm⁻³ on a 1-minute average with very high mass resolution (up to 14,000) ³¹. This
136 technique enhances our ability to explore the partitioning between the gas and particle phases
137 driving the aerosol behavior. In addition, during CLOUD-11, H₂SO₄ and oxidized organic
138 vapor concentrations were routinely measured with a nitrate-ion based Chemical Ionization
139 Atmospheric Pressure interface Time-of-Flight mass spectrometer (nitrate-Cl-APi-ToF) ⁷.
140 Calibration and corrections of sulfuric acid are similar to our previous work ³².

141 **Particle size distribution.** Particle size distributions were measured by several sizing
142 instruments optimized for different size ranges. Particles in the size range between 1.8–8 nm
143 were measured by the so-called DMA-train ⁹, and the corresponding growth-rate calculations
144 were performed with the appearance time method, which is effective in chamber experiments
145 where a clear front of a growing particle population can be identified during most nucleation
146 and growth events. Larger particles between 5–65 nm were measured with a scanning mobility
147 particle sizer (TSI nano-SMPS, model 3982).

148 **Volatility measurements.** Organic compounds with different vapor pressures have
149 distinct and reproducible thermograms during thermal desorption, and the maximum
150 desorption temperature (T_{\max}) correlates with the vaporization enthalpy and thus is a function
151 of the vapor pressure. We calibrated the thermal-desorption volatility measurements using a
152 suite of standards with known vapor pressures ^{30,33}. Briefly, we dissolved a set of organic acids
153 in isopropanol at low concentrations and manually deposited them on the FIGAERO filter with
154 a microliter syringe. We then held the mixture of acids at room temperature for 1 minute to
155 allow the solvent to evaporate before following the same thermal-desorption protocol used for
156 chamber experiments. We then established a thermal-desorption volatility calibration curve by

157 correlating $1/T_{\max}$ (1/K) with the sub-cooled liquid vapor pressure of each calibrant (Figure
158 S1). One potential source of uncertainty is that dimers may decompose to monomers when
159 heated, appearing as doublet-peaks in the thermogram³⁴. However, in this study, as dimer
160 formation is suppressed by NO_x, thermal decomposition of dimers did not affect the T_{\max} of
161 monomers.

162

163 RESULTS AND DISCUSSION

164 **Direct measurement of particle-phase volatility.** In the experiments described above
165 we directly measured the volatility of particle-phase products formed from photo-oxidation of
166 representative aromatic precursors (toluene and naphthalene) via their thermal-desorption
167 profile using an I⁻ FIGAERO-CIMS (see MATERIALS AND METHODS). In Figure 1, we
168 display the results in the two-dimensional volatility basis set (2D-VBS)³⁵, with volatility
169 (saturation concentration, C^* , in $\mu\text{g m}^{-3}$) on the x -axis and the oxygenation (oxygen to carbon
170 ratio, O:C) on the y -axis. We choose C^* and O:C as two complementary coordinates because
171 they explicitly track properties that we can observe in experiments and in the atmosphere, and
172 that we need to know to accurately model organic aerosol behavior.

173 In Figure 1A, we show monomers (crosses) and dimers (dots) derived from photo-
174 oxidation of toluene (blue symbols) and naphthalene (gray symbols) with NO_x present. Almost
175 all of the naphthalene-derived dimers appear in the ULVOC range³⁶, but the toluene-derived
176 dimers span the LVOC and ELVOC ranges. The measured monomers mostly appear in the
177 LVOC range, with a minority of the naphthalene-derived monomers being ELVOCs. Red
178 dashed lines are carbon numbers isopleths (C_6 , C_{11} and C_{19}); these are derived from the original
179 2D-VBS volatility parameterization, developed before the role of autoxidation in atmospheric
180 chemistry was appreciated³⁷, which assumed that a 50-50 mixture of -OH and =O
181 functionalities dominated the oxidized organics³⁵. These contours coincide well with our
182 measured volatilities from aromatic precursors. Toluene-derived monomers and naphthalene-
183 derived dimers fall along the C_6 and C_{19} isopleths, respectively, and toluene-derived dimers
184 and naphthalene-derived monomers overlap along the C_{11} isopleths. Measured monomers are
185 slightly more spread than dimers, potentially due to thermal decomposition or dehydration.

186 Similar to our aromatic oxidation experiments, in Figure 1B, we present monomers (green

crosses) and dimers (green dots) from α -pinene ozonolysis in the presence of NO_x . The carbon number isopleths (C_{10} and C_{19} , black dashed lines) are calculated using a revised VBS parameterization to reflect the contribution of autoxidation and subsequent –OOH functionality formation to initial particle growth ⁹. The measured dimers align with the C_{19} isopleth with great fidelity, though monomers are systematically but only slightly less volatile than the C_{10} isopleth. In this case, NO terminates the nascent peroxy radicals, suppresses autoxidation, and thus reduces the product oxygen numbers. As a result, the overall product volatility distribution is shifted towards the volatile end of the distribution compared to NO_x -free products (the oxygens on the $-\text{NO}_2$ moiety are not counted in O:C). Nevertheless, the products remain well described by the modified C^* vs O:C isopleths, with much higher volatility than the "traditional" VBS.

Compared with the α -pinene-derived oxidation products (C_{10}), toluene-derived products reach a similar volatility and O:C range with a C_7 backbone. More interestingly, naphthalene-derived oxidation products are much less volatile than α -pinene-derived products, though both start with a C_{10} backbone. The general volatility difference between dimers from these two precursors is ~5 orders of magnitude, as illustrated in Figure 1C and Figure S2. Aromatic oxidation appears to use oxygen atoms much more efficiently to form low-volatility products, consistent with the observation that bulk aromatic aerosols have a larger volume fraction remaining (VFR) after heating than biogenic aerosols ³⁸.

In Figure S3, we show the measured volatilities of aromatic- and α -pinene-derived oxidation products versus corresponding volatilities estimated using the VBS parametrizations in Donahue *et al.* ³⁵ and Stolzenburg *et al.* ⁹, respectively. Here, we choose organic compounds with six or more oxygen atoms to minimize the uncertainty of thermal decomposition in the FIGAERO. A good correlation (r^2 of 0.77) verifies the consistency between direct FIGAERO volatility measurements and VBS parameterizations.

Correlation between volatility and growth rate. Because the volatility of photo-oxidation products can vary by more than 10 orders of magnitude, it is convenient to simplify consideration of gas-particle partitioning by grouping compounds together within the VBS. In Figure 2A and B we show the binned volatility distribution of all organic particle-phase compounds observed with our FIGAERO-CIMS for the two representative aromatic

hydrocarbons. The volatility of toluene-derived oxidation products (Figure 2A) ranges from 10^{-9} to $10^{-2} \mu\text{g m}^{-3}$ with one mode at 10^{-4} and the other at $10^{-8} \mu\text{g m}^{-3}$. Naphthalene oxidation products are even less volatile (Figure 2B), with saturation concentrations ranging from 10^{-12} to $10^{-2} \mu\text{g m}^{-3}$. We divide the measured oxidation products into four groups: organonitrate monomers (light blue), pure organic monomers (light green), organonitrate dimers (navy blue) and pure organic dimers (dark green). For both of the monomers and dimers, organonitrates tend to be more volatile than non-nitrate products. This is again consistent with the NO_x effect on monoterpene oxidation and can also be explained either by the termination of nascent peroxy radicals or by reduced product reactivity with OH radicals and thus less secondary oxidation.

The measured volatility distributions are consistent with the 3–8 nm growth-rate measurements, where in Figure 2C we compare observed growth rates with gas-phase oxidation products measured with a nitrate-Cl-APi-TOF¹³. With most of the products in the LVOC range, toluene-derived oxidation products are somewhat less efficient than H_2SO_4 at driving growth. Naphthalene oxidation, on the other hand, produces a large amount of ULVOCs and ELVOCs, which are relatively more efficient at driving growth. The consistency between volatility and growth-rate measurements makes sense since all the species that can be measured in the particle phase must contribute to the particle growth, thus their volatilities can be taken as the criteria of growing particles and be used by VBS microphysical growth models.

The relative contribution of different precursors to new-particle formation depends on the volatility distribution of the products as well as the relative concentrations of the precursors. In polluted urban atmospheres, the concentration of aromatic hydrocarbons will exceed that of monoterpenes to be the dominant VOC precursors. Further, NO_x will substantially suppress the formation of monoterpene-derived ELVOCs, making the role of aromatic-derived oxidation products even more important. It is thus essential to add the volatility distribution of aromatic-derived oxidation products into particle-growth models and to assess the growth rates observed in urban areas.

Influence of functionality on volatility. Volatility is affected by both the molar weight and polarity of a molecule. Therefore, molecular structure and functionality matter. As a hydrocarbon chain gets larger, the increase in number of electrons increases the temporary

dipoles causing stronger van der Waals' forces and thus reducing the volatility. According to the SIMPOL model³⁹, each carbon atom lowers $\log C^*$ by 0.44. However, toluene has fewer carbon atoms than α -pinene, yet forms oxidation products with a comparable volatility distribution (see Figure 1). This is not due directly to aromaticity; each aromatic ring only lowers $\log C^*$ by 0.68. Even more notably, the 4 to 5 decades general difference between naphthalene-derived and α -pinene-derived oxidation products cannot be explained by different structures of their carbon backbones. This leaves functionality as the driving factor; the identity of the added functional groups plays a major role in volatility. Addition of one oxygen-containing functional group, in general, has a greater effect on volatility than a one-carbon increase in the size of the carbon skeleton, but the exact nature of the functional group is very important. Polar functional groups promote dipole-dipole interactions and thus will result in a larger decrease in volatility (e.g. each aldehyde group reduces $\log C^*$ by 1.35, and ketone group by 0.94). Functional groups that enable hydrogen bonding between molecules result in an even larger decrease in volatility (e.g. each hydroxyl group reduces $\log C^*$ by 2.23, hydroperoxide group by 2.48 (1.24 per oxygen), and carboxylic acid group by 3.58 (1.79 per oxygen))³⁹.

We investigate the influence of oxygen-containing functional groups on volatility of aromatic- and α -pinene-derived oxidation products by performing a parameterization using FIGAERO volatility data. This parameterization is based on a simple, three-parameter group-contribution expression:

$$\log_{10}C_{300}^* = (n_C^0 - n_C)b_C - n_O b_O \quad (1)$$

where C_{300}^* (at 300 K) is calculated from SIMPOL or directly measured by FIGAERO; $n_C^0 = 25$ is the carbon number n_C with $C^0 = 1 \mu\text{g m}^{-3}$ for pure hydrocarbons; $b_C = 0.48$ is the average effect of each added carbon on $\log_{10} C^*$; b_O indicates the average effect of each added oxygen on $\log_{10} C^*$ ³⁵. Note that a nitrate group ($-\text{ONO}_2$) typically reduces vapor pressure by about 2.5 orders of magnitude, so we replace any $-\text{ONO}_2$ group by $-\text{OH}$ for simplicity.

Here, instead of using this parameterization to calculate volatility, we fit the parameter b_O using volatility measured with FIGAERO thermal desorption. We performed the parameterization for products from α -pinene ozonolysis, toluene photo-oxidation and

276 naphthalene photo-oxidation, all with NO_x (green, blue and gray symbols in Figure 1,
277 respectively).

278 The fitted b_O value for α -pinene-derived oxidation products with NO_x present is 0.91. This
279 is consistent with hydroperoxide, ketone and aldehyde functionalities (roughly 1 decade per
280 oxygen) proposed in Tröstl *et al.*¹. To the extent that these simple group-contribution values
281 can be extrapolated to multifunctional compounds, this appears to be general within the
282 confines of this dataset; it certainly applies at the low-O:C limit, and it is reasonable for bi-
283 functional compounds (e.g. diacids⁴⁰). Many of the low-volatility particle-phase products from
284 monoterpene oxidation may thus be first-generation autoxidation products containing multiple
285 –OOH functional groups.

286 The fitted b_O values for toluene-derived and naphthalene-derived oxidation products, on
287 the other hand, are 1.55 and 1.72, respectively, suggesting more hydroxyl or carboxylic acid
288 functionality in aromatic-derived products. Oxygen-containing groups are very likely the key
289 to the systematic volatility difference between aromatic and monoterpene oxidation systems,
290 so we investigate the chemical mechanism in the next section in order to further understand the
291 formation of these functional groups in aromatic oxidation systems.

292 **Formation of oxidation products of aromatic hydrocarbons.** In this section, we use
293 toluene as the example to consider the chemical mechanism of aromatic oxidation. With this
294 proposed mechanism, we are able to address the volatility distributions arising from aromatics
295 and evaluate the contribution of aromatic-derived oxidation products to new-particle formation.

296 The detailed chemical mechanism of toluene oxidation in the atmosphere remains
297 uncertain. Wu *et al.*⁴¹ suggested that over 75% bicyclic peroxy radicals (BPRs) are formed
298 after the first OH oxidation of toluene based on a modified MCM mechanism. Wang *et al.*⁴²,
299 combining quantum calculations and flow-tube experiments, proposed that unimolecular H-
300 migration followed by O₂-addition, a so-called autoxidation step, can take place in BPRs,
301 which are important intermediates of the OH-initiated oxidation of alkyl benzenes⁴³. On the
302 other hand, Ji *et al.*⁴⁴ reported that cresols are much more stable than their corresponding
303 peroxy radicals⁴⁵, and, for the most favorable OH (ortho) addition, the pathway of H
304 abstraction by O₂ to form the cresol is dominant and O₂ addition to form the peroxy radical is
305 negligible. Other comprehensive studies^{46,47} identified ring-retaining products from cresol

306 oxidation and quantified both the yields of dihydroxy toluene and trihydroxy toluene being ~
307 0.7, which supported the results from Ji *et al.*⁴⁴.

308 These seemingly disparate findings may not be contradictory, since product formation
309 pathways and their yields depend strongly on experimental conditions. In the case of toluene
310 oxidation, the concentrations of OH radicals and O₂ control the relative importance of OH
311 oxidation and autoxidation, respectively. Wang *et al.*⁴² worked with [OH] of $(2.4\text{--}53) \times 10^4$
312 cm⁻³ in their experiments, which is one to two orders of magnitude lower than in ambient
313 conditions. The predominance of BPRs observed in their experiments could be due to
314 suppressed OH oxidation. However, in Ji's experiments⁴⁴, the concentration of O₂ was 10^{15}
315 cm⁻³, roughly 5000 times lower than in the atmospheric boundary layer ($[O_2] = 5 \times 10^{18}$ cm⁻³),
316 which could substantially suppress the formation of peroxy radicals.

317 In the CLOUD experiments, we set the experimental conditions as close as possible to
318 ambient conditions in order to acquire insights into the mechanisms that are atmospherically
319 relevant. We used [OH] of $\sim 10^6$ cm⁻³ and the [NO_x] of ~ 15 ppbv for the experiments in this
320 study. Also, the iodide reagent ion in our FIGAERO is sensitive to a wide range of species,
321 enabling us to observe oxygenated organics with oxygen numbers down to one or two with
322 good fidelity⁴⁸. Under these conditions, we observed products from both categories. In Figure
323 3 we show a mass defect plot of monomers from toluene oxidation measured in the gas phase.
324 The x-axis is the exact mass of the products and the y-axis is the mass defect. Each straight line
325 represents a homologous sequence of compounds with the same number of carbon, hydrogen,
326 and nitrogen atoms, but a different number of oxygen atoms. Notably, there are two main
327 categories of products: C₇H₈O₂₋₆, which likely preserves the aromaticity, and C₇H₁₀O₄₋₈, which
328 breaks the aromatic ring. These aromatic ring-retaining and aromatic ring-opening products are
329 associated with organonitrates, C₇H₇NO₃₋₇ and C₇H₉NO₅₋₁₀, respectively, via the termination
330 of the corresponding peroxy radicals by NO.

331 The detailed toluene oxidation mechanism is instructive (see Figure 4). Briefly, the rate of
332 OH addition to an aromatic ring is proportional to the electrophilic nature of the substituents
333 around the ring. Being electrophilic, the added OH groups in turn facilitate further addition
334 reactions by increasing the electron density on the ring through a resonance donating effect and
335 thus activating the aromatic ring. Following the ring-retaining pathway for toluene, OH radical

addition to the aromatic ring can occur multiple times to form cresol, dihydroxy-, trihydroxy- and even multi-OH substituted-toluene (i.e. $C_7H_8O_{2-6}$). The fact that we did not observe a significant amount of cresol could be due to its relatively slow production and fast consumption, which yields a low steady-state vapor concentration. This is consistent with the large signal we observed for $C_7H_7NO_3$, most likely hydroxy nitrotoluene, formed from the reaction of methyl phenoxy radicals with NO_2 or the reaction of cresol with NO_3 radicals⁴⁷.

Along the ring-opening pathway, peroxy radical production is followed by formation of bicyclic peroxides (i.e. $C_7H_{10}O_{4-8}$), which generates oxidation products with oxygen numbers of 8 or more. A succession of unimolecular H-migration reactions followed by O_2 addition (so-called autoxidation) can take place rapidly in bicyclic peroxy radicals (BPRs), which are important intermediates of the OH-initiated oxidation of aromatic compounds, possibly leading to the formation of highly oxygenated organic molecules (HOMs). However, with NO_x present in the chamber, peroxy-radical autoxidation is substantially suppressed, and we observed low HOM concentrations. Further, we also observed various decomposition products (C_5 and C_6) attributable to the bicyclic intermediate pathway. Although we distinguish the OH oxidation pathway from the autoxidation pathway, product formation in reality does not follow only one pathway. OH oxidation and autoxidation both can occur; it is their branching ratio under different conditions that matters.

Our purpose here is not to comprehensively define the detailed mechanism, but rather to demonstrate that aromatics are qualitatively different from monoterpenes in terms of the chemical mechanisms of photo-oxidation leading to new-particle formation and particle growth. From the proposed mechanism and from our FIGAERO observations, we conclude that rapid progression through multiple generations of OH oxidation is more pronounced in toluene and naphthalene than that in α -pinene, resulting in not only more oxidation but crucially also functional groups with much lower volatility per added oxygen atom, such as hydroxyl and carboxylic acids. For example, our proposed $C_7H_xO_{5-7}$ products from toluene contain up to 4 hydroxyl groups, while $C_{10}H_xO_{5-7}$, from α -pinene proposed in Tröstl *et al.*¹, only contain at most 1 hydroxyl group. Since hydroxyl groups have a much stronger effect per oxygen on volatility (stronger by about 1 decade per oxygen than $-OOH$), the larger number of hydroxyl groups accounts for the 4 to 5 decades difference in volatility between aromatic- and α -pinene-

366 derived oxidation products we observe in Figure 1.

367 These findings are consistent with findings for bulk SOA formation from monoterpenes
368 and aromatics⁴⁹. Specifically, with monoterpenes (especially for ozonolysis), SOA mass yields
369 can be reasonably described by first-generation products and a relatively stable product
370 distribution following the initial ozonolysis. For toluene photo-oxidation, on the other hand,
371 models of SOA formation in chambers must consider multiple generations of OH reaction
372 because the products become progressively more reactive with OH as the generation number
373 increases. To account for these differences, two different 2D-VBS representations of SOA
374 formation and aging have been developed for chemical transport models⁵⁰. Our findings
375 confirm this. The volatility parameter $b_0 \sim 1.6$ we derive for aromatic oxidation is consistent
376 with the original 2D-VBS parameterization representing aging and oxygenation of organic
377 products principally with an equal mixture of =O and –OH functional groups (and –C(O)OH).
378 This is a “traditional aging” scenario. On the other hand, the volatility parameter $b_0 \sim 0.9$ we
379 derive for α -pinene oxidation is consistent with the modified parameterizations more recently
380 employed to represent prompt formation of multiple –OOH functional groups. This is an
381 “autoxidation” scenario.

382 It is important to note that these scenarios are not mutually exclusive. As we discussed,
383 autoxidation does occur in the aromatic systems, and most of the first-generation monoterpene
384 products are not HOMs that have undergone autoxidation. All of these (mostly gas-phase)
385 products will almost certainly undergo further oxidation in the atmosphere, often on short
386 timescales⁵¹, and some of these second and later generation products will also undergo
387 autoxidation, as previous CLOUD experiments have shown for the first-generation
388 monoterpene product pinanediol^{16,52}.

389 Photo-oxidation of aromatic compounds (toluene and naphthalene) efficiently produces
390 low-volatility products that participate in new-particle formation and growth under conditions
391 typical of polluted urban settings. Here we have measured the composition of newly formed
392 particles with a FIGAERO-CIMS and directly constrained the volatility distribution of the
393 oxidation products that drive new-particle formation and growth. We find that oxidation of
394 toluene and naphthalene by OH radicals makes efficient use of added oxygen to form low-
395 volatility products, with much larger reductions in volatility per added oxygen atom than

396 corresponding α -pinene systems. This is consistent with known differences in the oxidation
397 mechanisms of aromatics and monoterpenes. Particle formation and growth in urban airmasses
398 may thus be strongly influenced by aromatic precursors.

399

400 ASSOCIATED CONTENT

401 Supporting Information

402 The Supporting Information is available free of charge on the ACS Publications website. This
403 file includes correlation of inverse maximum desorption temperature ($1/T_{\max}$) and the saturation
404 concentration ($\log C^*$) (Figure S1); example of thermograms (Figure S2); measured and
405 predicted volatility (Figure S3).

406

407 AUTHOR INFORMATION

408 Corresponding Author

409 *Neil M. Donahue (E-mail: nmd@andrew.cmu.edu).

410 Author Contributions

411 M. W., D. C., M. X., J. K., D. R. W., U. B., J. Dommen, I. E.-H., and N. M. D. designed
412 research; M. W., D. C., M. X., Q. Y., D. S., V. H., P. Y., A. L. V., R. L. M., A. A., A. B., B.
413 B., S. B., L. D., A. D., J. Duplissy, H. F., O. G., X. H., C. R. H., C. K., A. K., K. L., F. L., U.
414 M., T. P., V. P., L. L. J. Q., M. R., M. S., C. T., A. T., A. C. W., L. W., J. K., J. Dommen, and
415 I. E.-H. performed research; M. W., M. X., D. S., and N. M. D. contributed new
416 reagents/analytic tools; M. W., M. X., D. S., P. Y., and M. S. analyzed data; and M. W., D. C.,
417 M. X., A. L. V., U. M., J. K., U. B., J. Dommen, I. E.-H., and N. M. D. wrote the paper.

418 Notes

419 The authors declare no competing financial interest.

420

421 ACKNOWLEDGEMENT

422 We thank the European Organization for Nuclear Research (CERN) for supporting CLOUD
423 with important technical and financial resources and for providing a particle beam from the
424 CERN Proton Synchrotron. This research has received funding from the U.S. National Science
425 Foundation under grants AGS-1447056, AGS-1439551, AGS-1649147, AGS-1602086, and

426 AGS-1801897; the German Federal Ministry of Education and Research (No. 01LK1601A);
427 ERC-Consolidator Grant NANODYNAMITE 616075; Horizon 2020 Marie Skłodowska-
428 Curie Grant 656994 (“Nano-CAVa”); ERC Advanced “ATM-GP” grant No. 227463; the
429 Presidium of the Russian Academy of Sciences, the Program “High energy physics and
430 neutrino astrophysics” 2015; the Swiss National Science Foundation Projects 200020_152907,
431 20FI20_159851, 200021_169090, 200020_172602 and 20FI20_172622. The FIGAERO-
432 CIMS was supported by an MRI grant for the U.S. NSF AGS-1531284 as well as the Wallace
433 Research Foundation. O.G. thanks Doctoral Programme in Atmospheric Sciences at the
434 University of Helsinki for financial support.

435

436 **REFERENCES**

- 437 (1) Tröstl, J.; Chuang, W. K.; Gordon, H.; Heinritzi, M.; Yan, C.; Molteni, U.; Ahlm, L.; Frege, C.; Bianchi,
438 F.; Wagner, R.; Simon, M.; Lehtipalo, K.; Williamson, C.; Craven, J. S.; Duplissy, J.; Adamov, A.;
439 Almeida, J.; Bernhammer, A. K.; Breitenlechner, M.; Brilke, S.; Dias, A.; Ehrhart, S.; Flagan, R. C.;
440 Franchin, A.; Fuchs, C.; Guida, R.; Gysel, M.; Hansel, A.; Hoyle, C. R.; Jokinen, T.; Junninen, H.;
441 Kangasluoma, J.; Keskinen, H.; Kim, J.; Krapf, M.; Kürten, A.; Laaksonen, A.; Lawler, M.; Leiminger,
442 M.; Mathot, S.; Möhler, O.; Nieminen, T.; Onnela, A.; Petäjä, T.; Piel, F. M.; Miettinen, P.; Rissanen, M.
443 P.; Rondo, L.; Sarnela, N.; Schobesberger, S.; Sengupta, K.; Sipilä, M.; Smith, J. N.; Steiner, G.; Tomè,
444 A.; Virtanen, A.; Wagner, A. C.; Weingartner, E.; Wimmer, D.; Winkler, P. M.; Ye, P.; Carslaw, K. S.;
445 Curtius, J.; Dommen, J.; Kirkby, J.; Kulmala, M.; Riipinen, I.; Worsnop, D. R.; Donahue, N. M.;
446 Baltensperger, U. The Role of Low-Volatility Organic Compounds in Initial Particle Growth in the
447 Atmosphere. *Nature* **2016**, 533 (7604), 527–531. <https://doi.org/10.1038/nature18271>.
- 448 (2) Apte, J. S.; Marshall, J. D.; Cohen, A. J.; Brauer, M. Addressing Global Mortality from Ambient PM2.5.
449 *Env. Sci Technol* **2015**, 49 (13), 8057–8066. <https://doi.org/10.1021/acs.est.5b01236>.
- 450 (3) Lelieveld, J.; Evans, J. S.; Fnais, M.; Giannadaki, D.; Pozzer, A. The Contribution of Outdoor Air
451 Pollution Sources to Premature Mortality on a Global Scale. *Nature* **2015**, 525 (7569), 367–371.
452 <https://doi.org/10.1038/nature15371>.
- 453 (4) Riipinen, I.; Yli-Juuti, T.; Pierce, J. R.; Petaja, T.; Worsnop, D. R.; Kulmala, M.; Donahue, N. M.; Petäjä,
454 T.; Worsnop, D. R.; Kulmala, M.; Donahue, N. M. The Contribution of Organics to Atmospheric
455 Nanoparticle Growth. *Nat. Geosci.* **2012**, 5 (7), 453–458. <https://doi.org/10.1038/ngeo1499>.

- 456 (5) Anttila, T.; Kerminen, V.-M. M.; Lehtinen, K. E. J. J. Parameterizing the Formation Rate of New Particles:
457 The Effect of Nuclei Self-Coagulation. *J. Aerosol Sci.* **2010**, *41* (7), 621–636.
458 <https://doi.org/10.1016/j.jaerosci.2010.04.008>.
- 459 (6) Solomon, S.; Manning, M.; Marquis, M.; Qin, D.; others. *Climate Change 2007-the Physical Science
460 Basis: Working Group I Contribution to the Fourth Assessment Report of the IPCC*; Cambridge university
461 press, 2007; Vol. 4.
- 462 (7) Kirkby, J.; Duplissy, J.; Sengupta, K.; Frege, C.; Gordon, H.; Williamson, C.; Heinritzi, M.; Simon, M.;
463 Yan, C.; Almeida, J.; Trostl, J.; Nieminen, T.; Ortega, I. K.; Wagner, R.; Adamov, A.; Amorim, A.;
464 Bernhammer, A. K.; Bianchi, F.; Breitenlechner, M.; Brilke, S.; Chen, X. M.; Craven, J.; Dias, A.; Ehrhart,
465 Flagan, R. C.; Franchin, A.; Fuchs, C.; Guida, R.; Hakala, J.; Hoyle, C. R.; Jokinen, T.; Junninen, H.;
466 Kangasluoma, J.; Kim, J.; Krapf, M.; Kurten, A.; Laaksonen, A.; Lehtipalo, K.; Makhmutov, V.; Mathot,
467 S.; Molteni, U.; Onnela, A.; Perakyla, O.; Piel, F.; Petaja, T.; Praplan, A. P.; Pringle, K.; Rap, A.; Richards,
468 N. A. D. D.; Riipinen, I.; Rissanen, M. P.; Rondo, L.; Sarnela, N.; Schobesberger, S.; Scott, C. E.; Seinfeld,
469 J. H.; Sipila, M.; Steiner, G.; Stozhkov, Y.; Stratmann, F.; Tome, A.; Virtanen, A.; Vogel, A. L.; Wagner,
470 A. C.; Wagner, P. E.; Weingartner, E.; Wimmer, D.; Winkler, P. M.; Ye, P. L.; Zhang, X.; Hansel, A.;
471 Dommen, J.; Donahue, N. M.; Worsnop, D. R.; Baltensperger, U.; Kulmala, M.; Carslaw, K. S.; Curtius,
472 J.; Tomé, A.; Virtanen, A.; Vogel, A. L.; Wagner, A. C.; Wagner, P. E.; Weingartner, E.; Wimmer, D.;
473 Winkler, P. M.; Ye, P. L.; Zhang, X.; Hansel, A.; Dommen, J.; Donahue, N. M.; Worsnop, D. R.;
474 Baltensperger, U.; Kulmala, M.; Carslaw, K. S.; Curtius, J. Ion-Induced Nucleation of Pure Biogenic
475 Particles. *Nature* **2016**, *533* (7604), 521–526. <https://doi.org/10.1038/nature17953>.
- 476 (8) Lehtipalo, K.; Yan, C.; Dada, L.; Bianchi, F.; Xiao, M.; Wagner, R.; Stolzenburg, D.; Ahonen, L. R.;
477 Amorim, A.; Baccarini, A.; Bauer, P. S.; Baumgartner, B.; Bergen, A.; Bernhammer, A.-K.;
478 Breitenlechner, M.; Brilke, S.; Buchholz, A.; Mazon, S. B.; Chen, D.; Chen, X.; Dias, A.; Dommen, J.;
479 Draper, D. C.; Duplissy, J.; Ehn, M.; Finkenzeller, H.; Fischer, L.; Frege, C.; Fuchs, C.; Garmash, O.;
480 Gordon, H.; Hakala, J.; He, X.; Heikkinen, L.; Heinritzi, M.; Helm, J. C.; Hofbauer, V.; Hoyle, C. R.;
481 Jokinen, T.; Kangasluoma, J.; Kerminen, V.-M.; Kim, C.; Kirkby, J.; Kontkanen, J.; Kürten, A.; Lawler,
482 M. J.; Mai, H.; Mathot, S.; Mauldin, R. L.; Molteni, U.; Nichman, L.; Nie, W.; Nieminen, T.; Ojdanic,
483 A.; Onnela, A.; Passananti, M.; Petäjä, T.; Piel, F.; Pospisilova, V.; Quéléver, L. L. J.; Rissanen, M. P.;
484 Rose, C.; Sarnela, N.; Schallhart, S.; Schuchmann, S.; Sengupta, K.; Simon, M.; Sipilä, M.; Tauber, C.;
485 Tomé, A.; Tröstl, J.; Väisänen, O.; Vogel, A. L.; Volkamer, R.; Wagner, A. C.; Wang, M.; Weitz, L.;

- 486 Wimmer, D.; Ye, P.; Ylirisniö, A.; Zha, Q.; Carslaw, K. S.; Curtius, J.; Donahue, N. M.; Flagan, R. C.;
487 Hansel, A.; Riipinen, I.; Virtanen, A.; Winkler, P. M.; Baltensperger, U.; Kulmala, M.; Worsnop, D. R.
488 Multicomponent New Particle Formation from Sulfuric Acid, Ammonia, and Biogenic Vapors. *Sci. Adv.*
489 **2018**, *4* (12), eaau5363. <https://doi.org/10.1126/sciadv.aau5363>.
- 490 (9) Stolzenburg, D.; Fischer, L.; Vogel, A. L.; Heinritzi, M.; Schervish, M.; Simon, M.; Wagner, A. C.; Dada,
491 L.; Ahonen, L. R.; Amorim, A.; Baccarini, A.; Bauer, P. S.; Baumgartner, B.; Bergen, A.; Bianchi, F.;
492 Breitenlechner, M.; Brilke, S.; Buenrostro Mazon, S.; Chen, D.; Dias, A.; Draper, D. C.; Duplissy, J.; El
493 Haddad, I.; Finkenzeller, H.; Frege, C.; Fuchs, C.; Garmash, O.; Gordon, H.; He, X.; Helm, J.; Hofbauer,
494 V.; Hoyle, C. R.; Kim, C.; Kirkby, J.; Kontkanen, J.; Kürten, A.; Lampilahti, J.; Lawler, M.; Lehtipalo,
495 K.; Leiminger, M.; Mai, H.; Mathot, S.; Mentler, B.; Molteni, U.; Nie, W.; Nieminen, T.; Nowak, J. B.;
496 Ojdanic, A.; Onnela, A.; Passananti, M.; Petäjä, T.; Quéléver, L. L. J.; Rissanen, M. P.; Sarnela, N.;
497 Schallhart, S.; Tauber, C.; Tomé, A.; Wagner, R.; Wang, M.; Weitz, L.; Wimmer, D.; Xiao, M.; Yan, C.;
498 Ye, P.; Zha, Q.; Baltensperger, U.; Curtius, J.; Dommen, J.; Flagan, R. C.; Kulmala, M.; Smith, J. N.;
499 Worsnop, D. R.; Hansel, A.; Donahue, N. M.; Winkler, P. M. Rapid Growth of Organic Aerosol
500 Nanoparticles over a Wide Tropospheric Temperature Range. *Proc. Natl. Acad. Sci.* **2018**, *115* (37), 9122–
501 9127. <https://doi.org/10.1073/pnas.1807604115>.
- 502 (10) Donahue, N. M.; Ortega, I. K.; Chuang, W.; Riipinen, I.; Riccobono, F.; Schobesberger, S.; Dommen, J.;
503 Baltensperger, U.; Kulmala, M.; Worsnop, D. R.; Vehkamaki, H. How Do Organic Vapors Contribute to
504 New-Particle Formation? *Faraday Discuss.* **2013**, *165*, 91–104. <https://doi.org/10.1039/c3fd00046j>.
- 505 (11) Jokinen, T.; Sipila, M.; Richters, S.; Kerminen, V. M.; Paasonen, P.; Stratmann, F.; Worsnop, D.; Kulmala,
506 M.; Ehn, M.; Herrmann, H.; Berndt, T. Rapid Autoxidation Forms Highly Oxidized RO₂ Radicals in the
507 Atmosphere. *Angew. Chemie-International Ed.* **2014**, *53*, 14596–14600.
508 <https://doi.org/10.1002/anie.201408566>.
- 509 (12) Ehn, M.; Thornton, J. A.; Kleist, E.; Sipila, M.; Junninen, H.; Pullinen, I.; Springer, M.; Rubach, F.;
510 Tillmann, R.; Lee, B.; Lopez-Hilfiker, F.; Andres, S.; Acir, I. H.; Rissanen, M.; Jokinen, T.;
511 Schobesberger, S.; Kangasluoma, J.; Kontkanen, J.; Nieminen, T.; Kurten, T.; Nielsen, L. B.; Jorgensen,
512 S.; Kjaergaard, H. G.; Canagaratna, M.; Maso, M. D.; Berndt, T.; Petaja, T.; Wahner, A.; Kerminen, V.
513 M.; Kulmala, M.; Worsnop, D. R.; Wildt, J.; Mentel, T. F.; Dal Maso, M.; Berndt, T.; Petaja, T.; Wahner,
514 A.; Kerminen, V. M.; Kulmala, M.; Worsnop, D. R.; Wildt, J.; Mentel, T. F. A Large Source of Low-
515 Volatility Secondary Organic Aerosol. *Nature* **2014**, *506* (7489), 476–479.

- 516 https://doi.org/10.1038/nature13032.
- 517 (13) Yan, C.; Nie, W.; Vogel, A. L.; Dada, L.; Lehtipalo, K.; Stolzenburg, D.; Wagner, R.; Rissanen, M. P.;
518 Xiao, M.; Ahonen, L.; Fischer, L.; Rose, C.; Bianchi, F.; Gordon, H.; Simon, M.; Heinritzi, M.; Garmash,
519 O.; Roldin, P.; Dias, A.; Ye, P.; Hofbauer, V.; Amorim, A.; Bauer, P. S.; Bergen, A.; Bernhammer, A.-
520 K.; Breitenlechner, M.; Brilke, S.; Buchholz, A.; Mazon, S. B.; Canagaratna, M. R.; Chen, X.; Ding, A.;
521 Dommen, J.; Draper, D. C.; Duplissy, J.; Frege, C.; Heyn, C.; Guida, R.; Hakala, J.; Heikkinen, L.; Hoyle,
522 C. R.; Jokinen, T.; Kangasluoma, J.; Kirkby, J.; Kontkanen, J.; Kürten, A.; Lawler, M. J.; Mai, H.; Mathot,
523 S.; Mauldin, R. L.; Molteni, U.; Nichman, L.; Nieminen, T.; Nowak, J.; Ojdanic, A.; Onnela, A.; Pajunoja,
524 A.; Petäjä, T.; Piel, F.; Quéléver, L. L. J.; Sarnela, N.; Schallhart, S.; Sengupta, K.; Sipilä, M.; Tomé, A.;
525 Tröstl, J.; Väisänen, O.; Wagner, A. C.; Ylisirniö, A.; Zha, Q.; Baltensperger, U.; Carslaw, K. S.; Curtius,
526 J.; Flagan, R. C.; Hansel, A.; Riipinen, I.; Smith, J. N.; Virtanen, A.; Winkler, P. M.; Donahue, N. M.;
527 Kerminen, V.-M.; Kulmala, M.; Ehn, M.; Worsnop, D. R. Size-Dependent Influence of NO_x on the
528 Growth Rates of Organic Aerosol Particles. *Sci. Adv.* **2020**, *6* (22). <https://doi.org/10.1126/sciadv.aay4945>.
- 529 (14) Rissanen, M. P.; Kurtén, T.; Sipilä, M.; Thornton, J. A.; Kangasluoma, J.; Sarnela, N.; Junninen, H.;
530 Jørgensen, S.; Schallhart, S.; Kajos, M. K.; Taipale, R.; Springer, M.; Mentel, T. F.; Ruuskanen, T.; Petäjä,
531 T.; Worsnop, D. R.; Kjaergaard, H. G.; Ehn, M.; Kurtén, T.; Sipilä, M.; Thornton, J. A.; Kangasluoma,
532 J.; Sarnela, N.; Junninen, H.; Jørgensen, S.; Schallhart, S.; Kajos, M. K.; Kurten, T.; Sipila, M.; Thornton,
533 J. A.; Kangasluoma, J.; Sarnela, N.; Junninen, H.; Jorgensen, S.; Schallhart, S.; Kajos, M. K.; Taipale, R.;
534 Springer, M.; Mentel, T. F.; Ruuskanen, T.; Petaja, T.; Worsnop, D. R.; Kjaergaard, H. G.; Ehn, M. The
535 Formation of Highly Oxidized Multifunctional Products in the Ozonolysis of Cyclohexene. *J. Am. Chem.
536 Soc.* **2014**, *136*, 15596–15606. <https://doi.org/10.1021/ja507146s>.
- 537 (15) Donahue, N. M.; Henry, K. M.; Mentel, T. F.; Kiendler-Scharr, A.; Spindler, C.; Bohn, B.; Brauers, T.;
538 Dorn, H. P.; Fuchs, H.; Tillmann, R.; Wahner, A.; Saathoff, H.; Naumann, K.-H. H. K. H.; Möhler, O.;
539 Leisner, T.; Müller, L.; Reinnig, M.-C. C. M. C.; Hoffmann, T.; Salo, K.; Hallquist, M.; Frosch, M.; Bilde,
540 M.; Tritscher, T.; Barmet, P.; Praplan, A. P.; DeCarlo, P. F.; Dommen, J.; Prévôt, A. S. H.; Baltensperger,
541 U. Aging of Biogenic Secondary Organic Aerosol via Gas-Phase OH Radical Reactions. *Proc. Natl. Acad.
542 Sci.* **2012**, *109* (34), 13503–13508. <https://doi.org/10.1073/pnas.1115186109>.
- 543 (16) Schobesberger, S.; Junninen, H.; Bianchi, F.; Lönn, G.; Ehn, M.; Lehtipalo, K.; Dommen, J.; Ehrhart, S.;
544 Ortega, I. K.; Franchin, A.; Nieminen, T.; Riccobono, F.; Hutterli, M.; Duplissy, J.; Almeida, J.; Amorim,
545 A.; Breitenlechner, M.; Downard, A. J.; Dunne, E. M.; Flagan, R. C.; Kajos, M.; Keskinen, H.; Kirkby,

- 546 J.; Kupc, A.; Kürten, A.; Kurtén, T.; Laaksonen, A.; Mathot, S.; Onnela, A.; Praplan, A. P.; Rondo, L.;
547 Santos, F. D.; Schallhart, S.; Schnitzhofer, R.; Sipilä, M.; Tomé, A.; Tsagkogeorgas, G.; Vehkamäki, H.;
548 Wimmer, D.; Baltensperger, U.; Carslaw, K. S.; Curtius, J.; Hansel, A.; Petäjä, T.; Kulmala, M.; Donahue,
549 N. M.; Worsnop, D. R.; Lonn, G.; Ehn, M.; Lehtipalo, K.; Dommen, J.; Ehrhart, S.; Ortega, I. K.; Franchin,
550 A.; Nieminen, T.; Riccobono, F.; Hutterli, M.; Duplissy, J.; Almeida, J.; Amorim, A.; Breitenlechner, M.;
551 Downard, A. J.; Dunne, E. M.; Flagan, R. C.; Kajos, M.; Keskinen, H.; Kirkby, J.; Kupc, A.; Kurten, A.;
552 Kurten, T.; Laaksonen, A.; Mathot, S.; Onnela, A.; Praplan, A. P.; Rondo, L.; Santos, F. D.; Schallhart,
553 S.; Schnitzhofer, R.; Sipila, M.; Tome, A.; Tsagkogeorgas, G.; Vehkamaki, H.; Wimmer, D.;
554 Baltensperger, U.; Carslaw, K. S.; Curtius, J.; Hansel, A.; Petaja, T.; Kulmala, M.; Donahue, N. M.;
555 Worsnop, D. R. Molecular Understanding of Atmospheric Particle Formation from Sulfuric Acid and
556 Large Oxidized Organic Molecules. *Proc. Natl. Acad. Sci.* **2013**, *110* (43), 17223–17228.
557 <https://doi.org/10.1073/pnas.1306973110>.
- 558 (17) Koppmann, R. *Volatile Organic Compounds in the Atmosphere*; Wiley Online Library, 2007.
- 559 (18) Cabrera-Perez, D.; Taraborrelli, D.; Sander, R.; Pozzer, A. Global Atmospheric Budget of Simple
560 Monocyclic Aromatic Compounds. *Atmos. Chem. Phys.* **2016**, *16* (11).
- 561 (19) Shah, S. D.; Ogunyoku, T. A.; Miller, J. W.; Cocker, D. R. On-Road Emission Rates of PAH and n-Alkane
562 Compounds from Heavy-Duty Diesel Vehicles. *Environ. Sci. Technol.* **2005**, *39* (14), 5276–5284.
- 563 (20) Conde, F. J.; Ayala, J. H.; Afonso, A. M.; González, V. Emissions of Polycyclic Aromatic Hydrocarbons
564 from Combustion of Agricultural and Sylvicultural Debris. *Atmos. Environ.* **2005**, *39* (35), 6654–6663.
- 565 (21) McDonald, J. D.; Zielinska, B.; Fujita, E. M.; Sagebiel, J. C.; Chow, J. C.; Watson, J. G. Emissions from
566 Charbroiling and Grilling of Chicken and Beef. *J. Air Waste Manage. Assoc.* **2003**, *53* (2), 185–194.
- 567 (22) Bloss, C.; Wagner, V.; Jenkin, M. E.; Volkamer, R.; Bloss, W. J.; Lee, J. D.; Heard, D. E.; Wirtz, K.;
568 Martin-Reviejo, M.; Rea, G.; Wenger, J. C.; Pilling, M. J. Development of a Detailed Chemical
569 Mechanism (MCMv3.1) for the Atmospheric Oxidation of Aromatic Hydrocarbons. *Atmos. Chem. Phys.*
570 **2005**, *5*, 641–664. <https://doi.org/DOI 10.5194/acp-5-641-2005>.
- 571 (23) Offenberg, J. H.; Lewis, C. W.; Lewandowski, M.; Jaoui, M.; Kleindienst, T. E.; Edney, E. O.
572 Contributions of Toluene and α -Pinene to SOA Formed in an Irradiated Toluene/ α -Pinene/NO_x/Air
573 Mixture: Comparison of Results Using ¹⁴C Content and SOA Organic Tracer Methods. *Environ. Sci.*
574 *Technol.* **2007**, *41* (11), 3972–3976. <https://doi.org/10.1021/es070089+>.
- 575 (24) Odum, J. R.; Jungkamp, T. P. W. W.; Griffin, R. J.; Flagan, R. C.; Seinfeld, J. H. The Atmospheric

- 576 Aerosol-Forming Potential of Whole Gasoline Vapor. *Science* (80-.). **1997**, *276* (5309), 96–99.
577 <https://doi.org/DOI 10.1126/science.276.5309.96>.
- 578 (25) Forstner, H. J. L.; Flagan, R. C.; Seinfeld, J. H. Secondary Organic Aerosol from the Photooxidation of
579 Aromatic Hydrocarbons: Molecular Composition. *Environ. Sci. Technol.* **1997**, *31* (5), 1345–1358.
580 <https://doi.org/10.1021/es9605376>.
- 581 (26) Molteni, U.; Bianchi, F.; Klein, F.; El Haddad, I.; Frege, C.; Rossi, M. J.; Dommen, J.; Baltensperger, U.
582 Formation of Highly Oxygenated Organic Molecules from Aromatic Compounds. *Atmos. Chem. Phys.*
583 **2018**, *18*, 1909–1921. <https://doi.org/10.5194/acp-18-1909-2018>.
- 584 (27) Garmash, O.; Rissanen, M. P.; Pullinen, I.; Schmitt, S.; Kausiala, O.; Tillmann, R.; Zhao, D.; Percival,
585 C.; Bannan, T. J.; Priestley, M.; Hallquist, A. M.; Kleist, E.; Kiendler-Scharr, A.; Hallquist, M.; Berndt,
586 T.; McFiggans, G.; Wildt, J.; Mentel, T. F.; Ehn, M. Multi-Generation OH Oxidation as a Source for
587 Highly Oxygenated Organic Molecules from Aromatics. *Atmos. Chem. Phys.* **2020**, *20*, 515–537.
588 <https://doi.org/10.5194/acp-20-515-2020>.
- 589 (28) Dias, A.; Ehrhart, S.; Vogel, A.; Mathot, S.; Onnela, A.; Almeida, J.; Kirkby, J.; Williamson, C.; Mumford,
590 S. Temperature Uniformity in the CERN CLOUD Chamber. *Atmos. Meas. Tech.* **2017**, *10*, 5075–5088.
- 591 (29) Bertram, T. H.; Kimmel, J. R.; Crisp, T. A.; Ryder, O. S.; Yatavelli, R. L. N.; Thornton, J. A.; Cubison,
592 M. J.; Gonin, M.; Worsnop, D. R. A Field-Deployable, Chemical Ionization Time-of-Flight Mass
593 Spectrometer. *Atmos. Meas. Tech.* **2011**, *4*, 1471–1479. <https://doi.org/10.5194/amt-4-1471-2011>.
- 594 (30) Lopez-Hilfiker, F. D.; Mohr, C.; Ehn, M.; Rubach, F.; Kleist, E.; Wildt, J.; Mentel, T. F.; Lutz, A.;
595 Hallquist, M.; Worsnop, D.; Thornton, J. A. A Novel Method for Online Analysis of Gas and Particle
596 Composition: Description and Evaluation of a Filter Inlet for Gases and AEROSols (FIGAERO). *Atmos.*
597 *Meas. Tech.* **2014**, *7* (4), 983–1001. <https://doi.org/10.5194/amt-7-983-2014>.
- 598 (31) Junninen, H.; Ehn, M.; Petaja, T.; Luosujarvi, L.; Kotiaho, T.; Kostainen, R.; Rohner, U.; Gonin, M.;
599 Fuhrer, K.; Kulmala, M.; Worsnop, D. R.; Petäjä, T.; Luosujärvi, L.; Kotiaho, T.; Kostainen, R.; Rohner,
600 U.; Gonin, M.; Fuhrer, K.; Kulmala, M.; Worsnop, D. R. A High-Resolution Mass Spectrometer to
601 Measure Atmospheric Ion Composition. *Atmos. Meas. Tech.* **2010**, *3* (4), 1039–1053.
602 <https://doi.org/10.5194/amt-3-1039-2010>.
- 603 (32) Kurten, A.; Jokinen, T.; Simon, M.; Sipila, M.; Sarnela, N.; Junninen, H.; Adamov, A.; Almeida, J.;
604 Amorim, A.; Bianchi, F.; Breitenlechner, M.; Dommen, J.; Donahue, N. M.; Duplissy, J.; Ehrhart, S.;
605 Flagan, R. C.; Franchin, A.; Hakala, J.; Hansel, A.; Heinritzi, M.; Hutterli, M.; Kangasluoma, J.; Kirkby,

- 606 J.; Laaksonen, A.; Lehtipalo, K.; Leiminger, M.; Makhmutov, V.; Mathot, S.; Onnela, A.; Petaja, T.;
607 Praplan, A. P.; Riccobono, F.; Rissanen, M. P.; Rondo, L.; Schobesberger, S.; Seinfeld, J. H.; Steiner, G.;
608 Tome, A.; Trostl, J.; Winkler, P. M.; Williamson, C.; Wimmer, D.; Ye, P. L.; Baltensperger, U.; Carslaw,
609 K. S.; Kulmala, M.; Worsnop, D. R.; Curtius, J.; Kurtén, A.; Jokinen, T.; Simon, M.; Sipilä, M.; Sarnela,
610 N.; Junninen, H.; Adamov, A.; Almeida, J.; Amorim, A.; Bianchi, F.; Breitenlechner, M.; Dommen, J.;
611 Donahue, N. M.; Duplissy, J.; Ehrhart, S.; Flagan, R. C.; Franchin, A.; Hakala, J.; Hansel, A.; Heinritzi,
612 M.; Hutterli, M.; Kangasluoma, J.; Kirkby, J.; Laaksonen, A.; Lehtipalo, K.; Leiminger, M.; Makhmutov,
613 V.; Mathot, S.; Onnela, A.; Petäjä, T.; Praplan, A. P.; Riccobono, F.; Rissanen, M. P.; Rondo, L.;
614 Schobesberger, S.; Seinfeld, J. H.; Steiner, G.; Tomé, A.; Tröstl, J.; Winkler, P. M.; Williamson, C.;
615 Wimmer, D.; Ye, P. L.; Baltensperger, U.; Carslaw, K. S.; Kulmala, M.; Worsnop, D. R.; Curtius, J.
616 Neutral Molecular Cluster Formation of Sulfuric Acid-Dimethylamine Observed in Real Time under
617 Atmospheric Conditions. *Proc. Natl. Acad. Sci. U. S. A.* **2014**, *111* (42), 15019–15024.
618 <https://doi.org/10.1073/pnas.1404853111>.
- 619 (33) Stark, H.; Yatavelli, R. L. N.; Thompson, S. L.; Kang, H.; Krechmer, J. E.; Kimmel, J. R.; Palm, B. B.;
620 Hu, W. W.; Hayes, P. L.; Day, D. A.; Campuzano-Jost, P.; Canagaratna, M. R.; Jayne, J. T.; Worsnop, D.
621 R.; Jimenez, J. L. Impact of Thermal Decomposition on Thermal Desorption Instruments: Advantage of
622 Thermogram Analysis for Quantifying Volatility Distributions of Organic Species. *Environ. Sci. Technol.*
623 **2017**, *51*, 8491–8500. <https://doi.org/10.1021/acs.est.7b00160>.
- 624 (34) Wang, M. Y.; Yao, L.; Zheng, J.; Wang, X. K.; Chen, J. M.; Yang, X.; Worsnop, D. R.; Donahue, N. M.;
625 Wang, L. Reactions of Atmospheric Particulate Stabilized Criegee Intermediates Lead to High-Molecular-
626 Weight Aerosol Components. *Environ. Sci. Technol.* **2016**, *50*, 5702–5710.
627 <https://doi.org/10.1021/acs.est.6b02114>.
- 628 (35) Donahue, N. M.; Epstein, S. A.; Pandis, S. N.; Robinson, A. L. A Two-Dimensional Volatility Basis Set:
629 1. Organic-Aerosol Mixing Thermodynamics. *Atmos. Chem. Phys.* **2011**, *11* (7), 3303–3318.
630 <https://doi.org/10.5194/acp-11-3303-2011>.
- 631 (36) Schervish, M.; Donahue, N. M. Peroxy Radical Chemistry and the Volatility Basis Set. *Atmos. Chem.*
632 *Phys.* **2020**, *20*, 1183–1199. <https://doi.org/10.5194/acp-20-1183-2020>.
- 633 (37) Bianchi, F.; Kurtén, T.; Riva, M.; Mohr, C.; Rissanen, M. P.; Roldin, P.; Berndt, T.; Crounse, J. D.;
634 Wennberg, P. O.; Mentel, T. F.; Wildt, J.; Junninen, H.; Jokinen, T.; Kulmala, M.; Worsnop, D. R.;
635 Thornton, J. A.; Donahue, N.; Kjaergaard, H. G.; Ehn, M. Highly Oxygenated Organic Molecules (HOM)

- 636 from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol.
- 637 *Chem. Rev.* **2019**, *119* (6), 3472–3509. <https://doi.org/10.1021/acs.chemrev.8b00395>.
- 638 (38) Emanuelsson, E. U.; Hallquist, M.; Kristensen, K.; Glasius, M.; Bohn, B.; Fuchs, H.; Kammer, B.;
639 Kiendler-Scharr, A.; Nehr, S.; Rubach, F.; Tillmann, R.; Wahner, A.; Wu, H.-C. C.; Mentel, T. F.
640 Formation of Anthropogenic Secondary Organic Aerosol (SOA) and Its Influence on Biogenic SOA
641 Properties. *Atmos. Chem. Phys.* **2013**, *13* (5), 2837–2855. <https://doi.org/10.5194/acp-13-2837-2013>.
- 642 (39) Pankow, J. F.; Asher, W. E. SIMPOL.1: A Simple Group Contribution Method for Predicting Vapor
643 Pressures and Enthalpies of Vaporization of Multifunctional Organic Compounds. *Atmos. Chem. Phys.*
644 **2008**, *8*, 2773–2796. <https://doi.org/DOI 10.5194/acp-8-2773-2008>.
- 645 (40) Bilde, M.; Barsanti, K.; Booth, M.; Cappa, C. D.; Donahue, N. M.; Emanuelsson, E. U.; McFiggans, G.;
646 Krieger, U. K.; Marcolli, C.; Topping, D.; Ziemann, P.; Barley, M.; Clegg, S.; Dennis-Smith, B.;
647 Hallquist, M.; Hallquist, Å. M.; Khlystov, A.; Kulmala, M.; Mogensen, D.; Percival, C. J.; Pope, F.; Reid,
648 J. P.; Ribeiro da Silva, M. A. V.; Rosenoern, T.; Salo, K.; Soonsin, V. P.; Yli-Juuti, T.; Prisle, N. L.; Pagels,
649 J.; Rarey, J.; Zardini, A. A.; Riipinen, I. Saturation Vapor Pressures and Transition Enthalpies of Low-
650 Volatility Organic Molecules of Atmospheric Relevance: From Dicarboxylic Acids to Complex Mixtures.
651 *Chem. Rev.* **2015**, *115* (10), 4115–4156. <https://doi.org/10.1021/cr5005502>.
- 652 (41) Wu, R. R.; Pan, S. S.; Li, Y.; Wang, L. M. Atmospheric Oxidation Mechanism of Toluene. *J. Phys. Chem. A*
653 **2014**, *118* (25), 4533–4547. <https://doi.org/10.1021/jp500077f>.
- 654 (42) Wang, S. N.; Wu, R. R.; Berndt, T.; Ehn, M.; Wang, L. M. Formation of Highly Oxidized Radicals and
655 Multifunctional Products from the Atmospheric Oxidation of Alkylbenzenes. *Environ. Sci. Technol.* **2017**,
656 *51*, 8442–8449. <https://doi.org/10.1021/acs.est.7b02374>.
- 657 (43) Volkamer, R.; Platt, U.; Wirtz, K. Primary and Secondary Glyoxal Formation from Aromatics:
658 Experimental Evidence for the Bicycloalkyl - Radical Pathway from Benzene, Toluene, and p-Xylene. *J. Phys. Chem. A*
659 **2001**, *105* (33), 7865–7874. <https://doi.org/10.1021/jp010152w>.
- 660 (44) Ji, Y. M.; Zhao, J.; Terazono, H.; Misawa, K.; Levitt, N. P.; Li, Y. X.; Lin, Y.; Peng, J. F.; Wang, Y.;
661 Duan, L.; Pan, B. W.; Zhang, F.; Feng, X. D.; An, T. C.; Marrero-Ortiz, W.; Secret, J.; Zhang, A. L.;
662 Shibuya, K.; Molina, M. J.; Zhang, R. Y. Reassessing the Atmospheric Oxidation Mechanism of Toluene.
663 *Proc. Natl. Acad. Sci. U. S. A.* **2017**, *114*, 8169–8174. <https://doi.org/10.1073/pnas.1705463114>.
- 664 (45) Klotz, B.; Sørensen, S.; Barnes, I.; Becker, K. H.; Etzkorn, T.; Volkamer, R.; Platt, U.; Wirtz, K.; Martin-
665 Reviejo, M. Atmospheric Oxidation of Toluene in a Large-Volume Outdoor Photoreactor: In Situ

- 666 Determination of Ring-Retaining Product Yields. *J. Phys. Chem. A* **1998**, *102* (50), 10289–10299.
- 667 <https://doi.org/10.1021/jp982719n>.
- 668 (46) Olariu, R. I.; Klotz, B.; Barnes, I.; Becker, K. H.; Mocanu, R. FT-IR Study of the Ring-Retaining Products
669 from the Reaction of OH Radicals with Phenol, o-, m-, and p-Cresol. *Atmos. Environ.* **2002**, *36* (22),
670 3685–3697. [https://doi.org/10.1016/S1352-2310\(02\)00202-9](https://doi.org/10.1016/S1352-2310(02)00202-9).
- 671 (47) Schwantes, R. H.; Schilling, K. A.; McVay, R. C.; Lignell, H.; Coggon, M. M.; Zhang, X.; Wennberg, P.
672 O.; Seinfeld, J. H. Formation of Highly Oxygenated Low-Volatility Products from Cresol Oxidation.
673 *Atmos. Chem. Phys.* **2017**, *17* (5), 3453–3474. <https://doi.org/10.5194/acp-17-3453-2017>.
- 674 (48) Iyer, S.; Lopez-Hilfiker, F.; Lee, B. H.; Thornton, J. A.; Kurtén, T.; Kurten, T. Modeling the Detection of
675 Organic and Inorganic Compounds Using Iodide-Based Chemical Ionization. *J. Phys. Chem. A* **2016**, *120*
676 (4), 576–587. <https://doi.org/10.1021/acs.jpca.5b09837>.
- 677 (49) Zhao, B.; Wang, S.; Donahue, N. M.; Chuang, W.; Hildebrandt Ruiz, L.; Ng, N. L.; Wang, Y.; Hao, J.
678 Evaluation of One-Dimensional and Two-Dimensional Volatility Basis Sets in Simulating the Aging of
679 Secondary Organic Aerosol with Smog-Chamber Experiments. *Environ. Sci. Technol.* **2015**, *49* (4), 2245–
680 2254. <https://doi.org/10.1021/es5048914>.
- 681 (50) Zhao, B.; Wang, S.; Donahue, N. M.; Jathar, S. H.; Huang, X.; Wu, W.; Hao, J.; Robinson, A. L.
682 Quantifying the Effect of Organic Aerosol Aging and Intermediate-Volatility Emissions on Regional-
683 Scale Aerosol Pollution in China. *Sci. Rep.* **2016**, *6* (June), 28815. <https://doi.org/10.1038/srep28815>.
- 684 (51) Donahue, N. M.; Chuang, W.; Epstein, S. A.; Kroll, J. H.; Worsnop, D. R.; Robinson, A. L.; Adams, P.
685 J.; Pandis, S. N. Why Do Organic Aerosols Exist? Understanding Aerosol Lifetimes Using the Two-
686 Dimensional Volatility Basis Set. *Environ. Chem.* **2013**, *10*, 151–157.
- 687 (52) Riccobono, F.; Schobesberger, S.; Scott, C. E.; Dommen, J.; Ortega, I. K.; Rondo, L.; Almeida, J.;
688 Amorim, A.; Bianchi, F.; Breitenlechner, M.; David, A.; Downard, A.; Dunne, E. M.; Duplissy, J.; Ehrhart,
689 S.; Flagan, R. C.; Franchin, A.; Hansel, A.; Junninen, H.; Kajos, M.; Keskinen, H.; Kupc, A.; Kurten, A.;
690 Kvashin, A. N.; Laaksonen, A.; Lehtipalo, K.; Makhmutov, V.; Mathot, S.; Nieminen, T.; Onnela, A.;
691 Petaja, T.; Praplan, A. P.; Santos, F. D.; Schallhart, S.; Seinfeld, J. H.; Sipila, M.; Spracklen, D. V.;
692 Stozhkov, Y.; Stratmann, F.; Tome, A.; Tsagkogeorgas, G.; Vaattovaara, P.; Viisanen, Y.; Vrtala, A.;
693 Wagner, P. E.; Weingartner, E.; Wex, H.; Wimmer, D.; Carslaw, K. S.; Curtius, J.; Donahue, N. M.;
694 Kirkby, J.; Kulmala, M.; Worsnop, D. R.; Baltensperger, U.; Kürten, A.; Kvashin, A. N.; Laaksonen, A.;
695 Lehtipalo, K.; Makhmutov, V.; Mathot, S.; Nieminen, T.; Onnela, A.; Petäjä, T.; Praplan, A. P.; Santos,

696 F. D.; Schallhart, S.; Seinfeld, J. H.; Sipilä, M.; Spracklen, D. V.; Stozhkov, Y.; Stratmann, F.; Tomé, A.;
697 Tsagkogeorgas, G.; Vaattovaara, P.; Viisanen, Y.; Vrtala, A.; Wagner, P. E.; Weingartner, E.; Wex, H.;
698 Wimmer, D.; Carslaw, K. S.; Curtius, J.; Donahue, N. M.; Kirkby, J.; Kulmala, M.; Worsnop, D. R.;
699 Baltensperger, U. Oxidation Products of Biogenic Emissions Contribute to Nucleation of Atmospheric
700 Particles. *Science (80-)*. **2014**, *344*, 717–721. <https://doi.org/10.1126/science.1243527>.

701 (53) Nieminen, T.; Lehtinen, K. E. J.; Kulmala, M. Sub-10 Nm Particle Growth by Vapor Condensation –
702 Effects of Vapor Molecule Size and Particle Thermal Speed. *Atmos. Chem. Phys.* **2010**, *10* (20), 9773–
703 9779. <https://doi.org/10.5194/acp-10-9773-2010>.

704

FIGURE LEGENDS

705
706
707 **Figure 1.** Directly measured particle-phase oxidation products plotted on a two-dimensional
708 VBS. The background colors indicate the saturation concentration (C^*) range of ultra-low
709 volatility (ULVOCs, purple), extremely low volatility (ELVOCs, gray), low volatility (LVOCs,
710 pink) and semi-volatile (SVOCs, green) organic compounds. (A) Monomers (crosses) and
711 dimers (dots) from toluene (blue symbols) and naphthalene (gray symbols) photo-oxidation
712 with NO_x present. The volatility distribution can be represented by carbon numbers isopleths
713 (red dashed lines) derived from the original volatility basis set (VBS) parameterization³⁵. (B)
714 Monomers (green crosses) and dimers (green dots) from α -pinene ozonolysis with NO_x present
715 observed in CLOUD 11 experiments as well as carbon numbers isopleths (black dashed lines)
716 calculated using a revised VBS parameterization to reflect the –OOH functionality formed via
717 autoxidation⁹. (C) Direct comparison of aromatic-derived dimers and α -pinene-derived dimers.
718 Both are C_{20} but the aromatics are much less volatile than α -pinene-derived dimers. Symbol
719 sizes are proportional to the particle-phase signals.

720

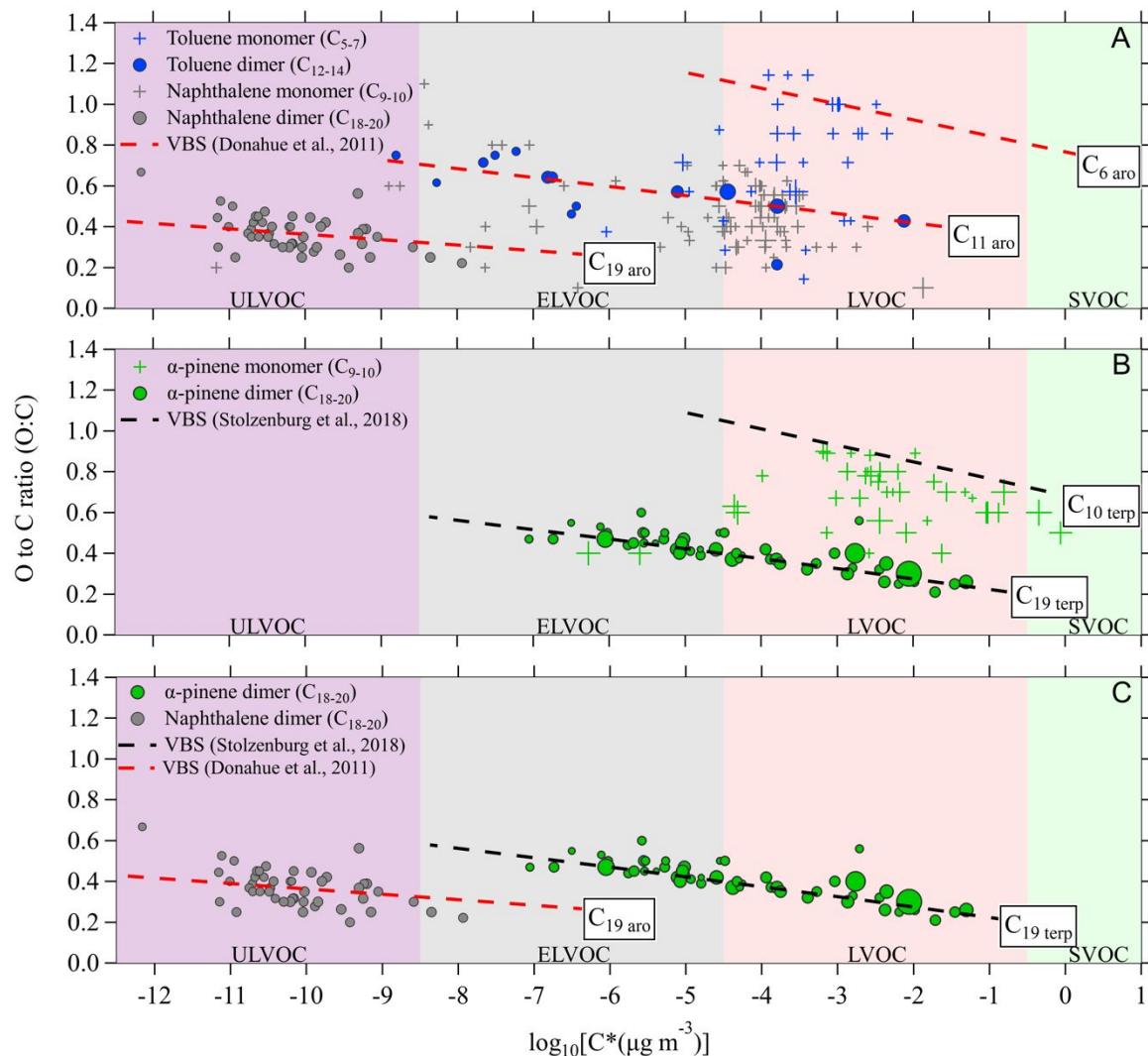
721 **Figure 2.** Directly measured volatility distributions of particle-phase products formed during
722 toluene (A) and naphthalene (B) photo-oxidation with NO_x present. Products are binned with
723 the measured relative mass concentration in percent. The measured oxidation products are
724 divided into organonitrate monomers (light blue), pure organic monomers (light green),
725 organonitrate dimers (navy blue) and pure organic dimers (dark green). (C) Growth rates
726 measured with DMA train vs the sum of sulfuric acid and HOM concentrations. Sulfuric acid
727 kinetic limit is taken from Nieminen *et al.*⁵³.

728

729 **Figure 3.** The mass defect plot showing products from toluene oxidation measured in the gas
730 phase. The x-axis is the exact mass of oxidation products and y-axis is the mass defect. The
731 color of circles denotes the type of products, and their size is proportional to the logarithm of
732 the count rate. Each straight line represents a group of compounds with the same number of
733 carbon, hydrogen, and nitrogen atoms, but a different number of oxygen atoms.

734

735 **Figure 4.** Gas-phase chemical mechanism for toluene photo-oxidation with NO_x present. The
736 reaction starts with OH addition to the aromatic ring, resulting in oxidized radicals. Closed-
737 shell products will be formed after radicals undergo various termination pathways. Products
738 outlined are detected by the iodide-adduct CIMS. Non-nitrate products and organonitrate
739 products are marked with green and blue, respectively. (Adapted from Schwantes *et al.*⁴⁷.)
740



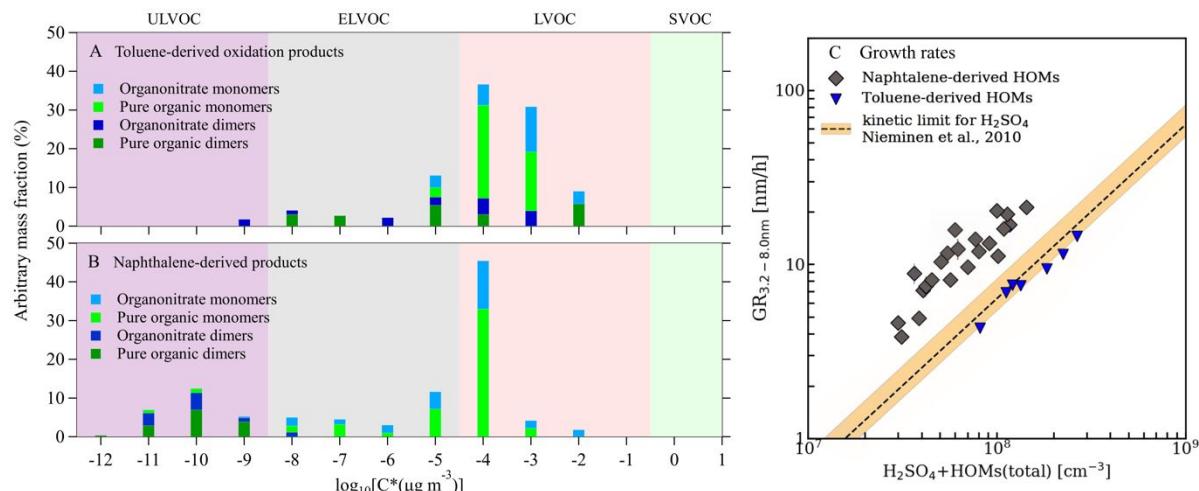
741

742

743

744

Figure 1.



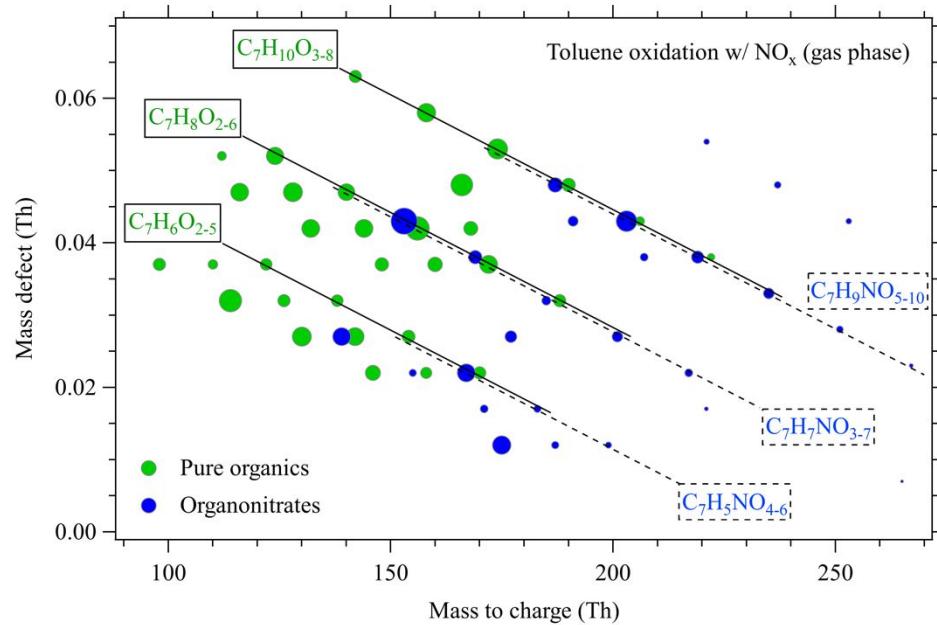
745

746

747

748

Figure 2.



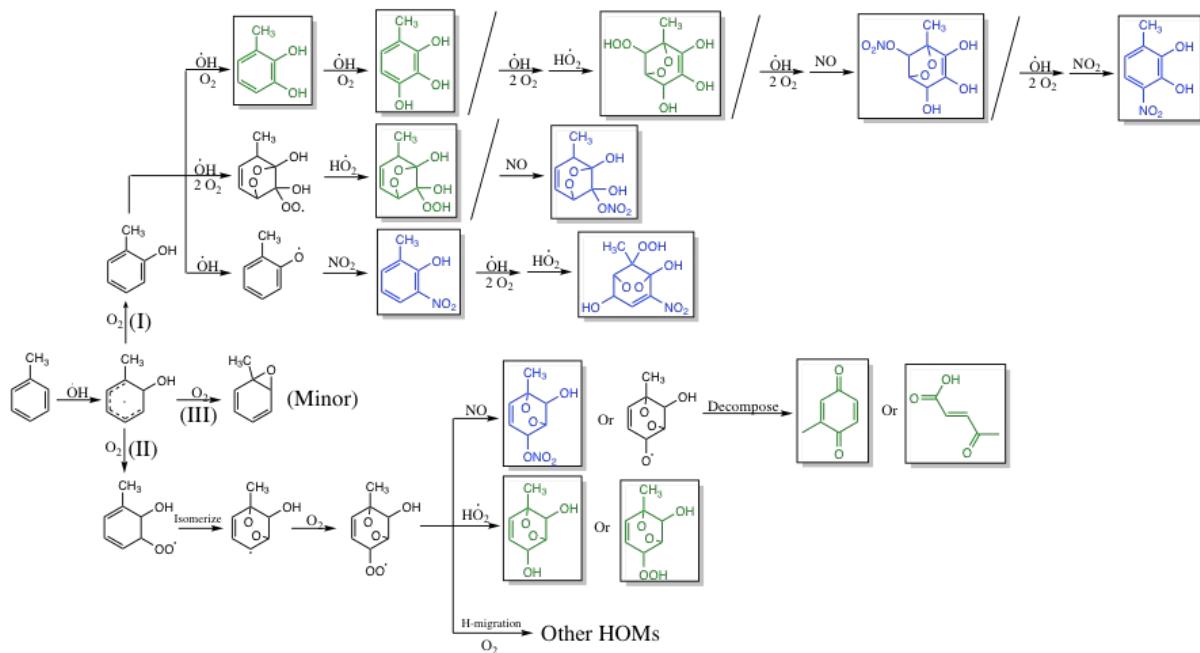
749

750

751

752

Figure 3.



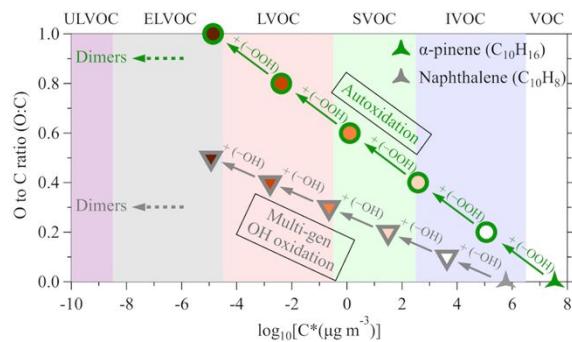
753

754

755

756

Figure 4.



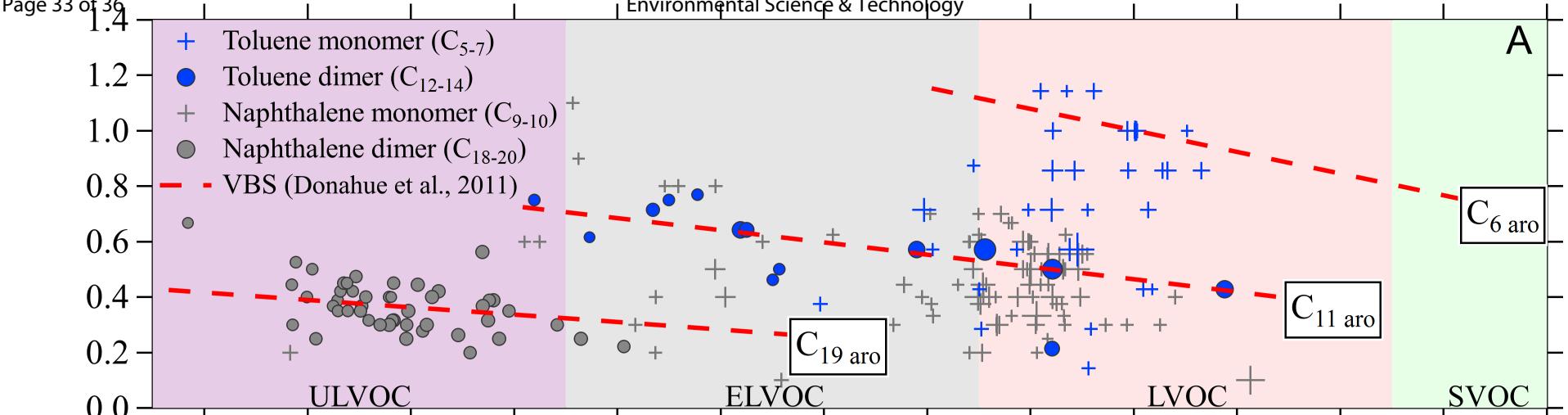
757

TOC art

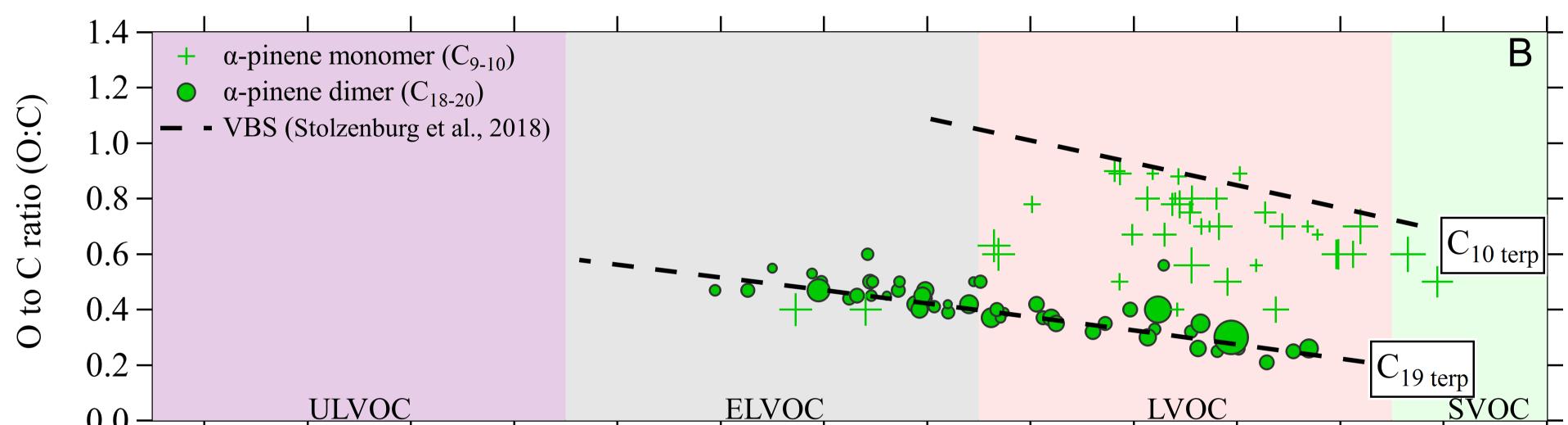
758

759

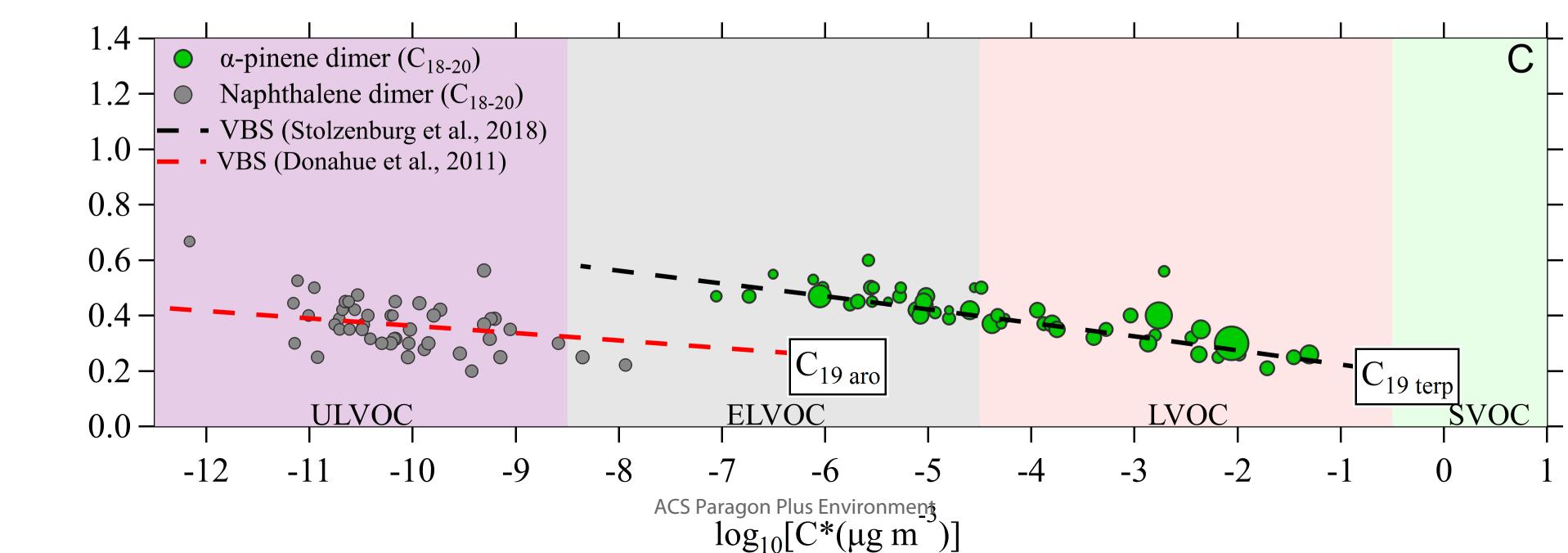
A

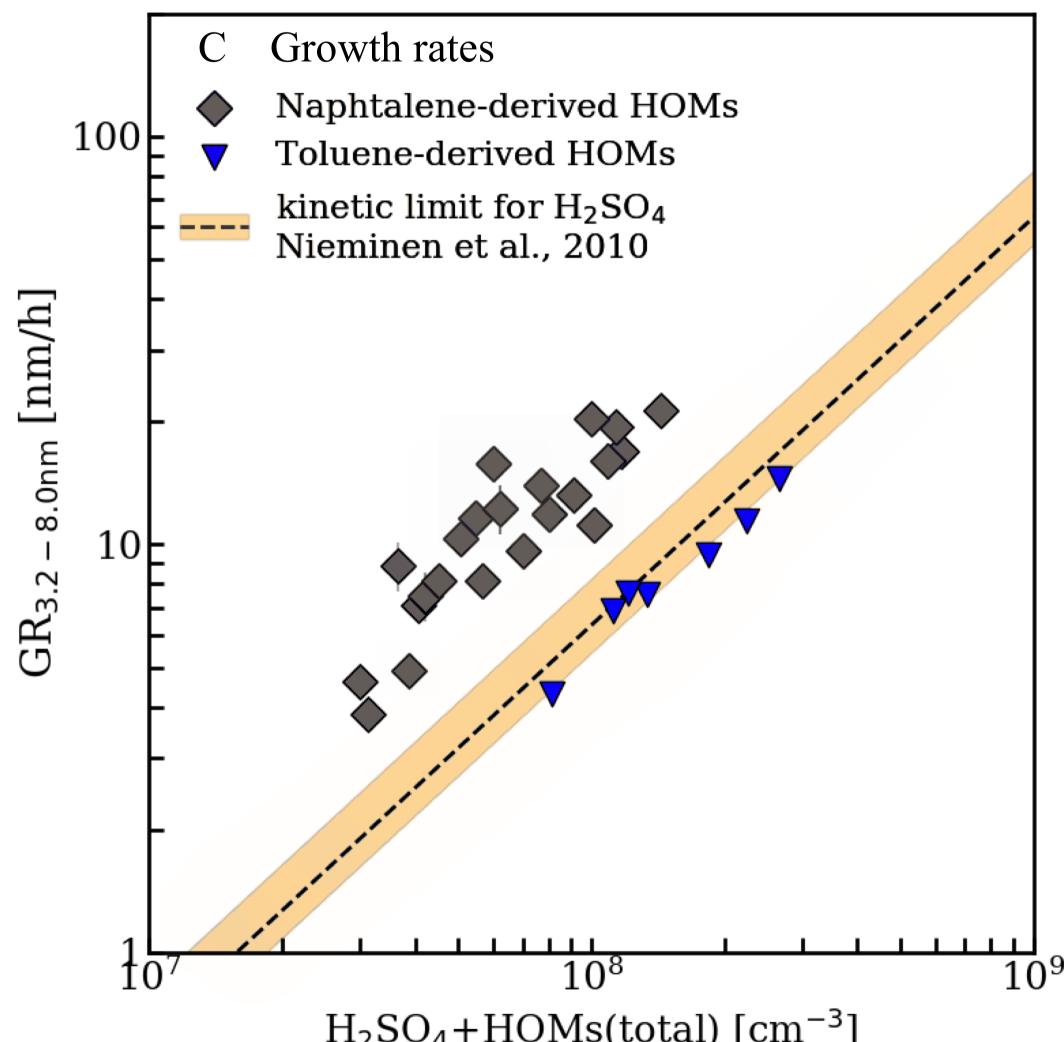
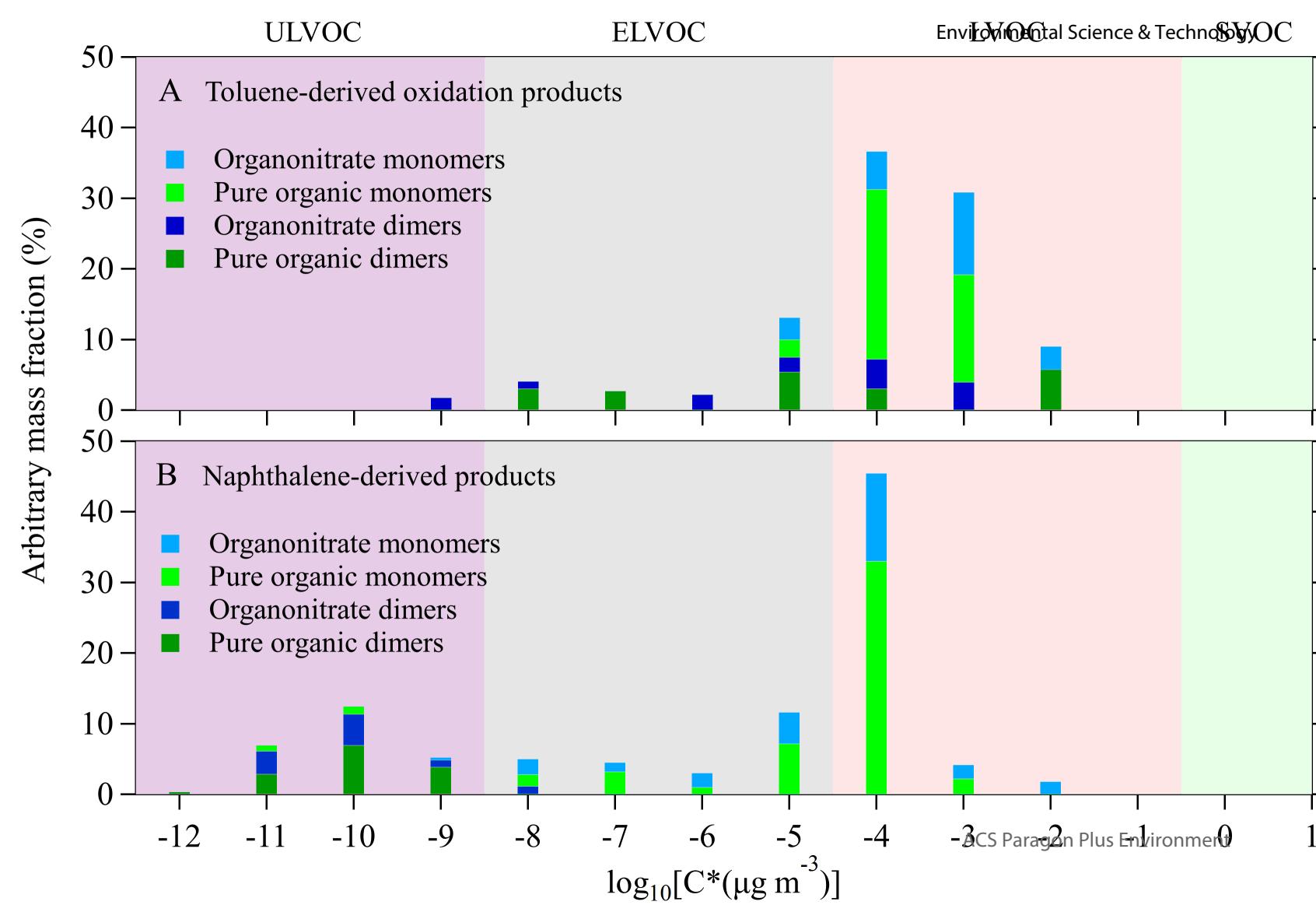


B



C





Toluene oxidation w/ NO_x (gas phase)

Mass defect (Th)

0.06

0.04

0.02

0.00



Pure organics

Organonitrates

100

150

200

250

ACS Paragon Plus Environment

Mass to charge (Th)

